

海流流速计算的研究

景振华

(山东海洋学院)

前 言

由于大范围同步连续观测海流流速很困难,这才产生建立一定的理论及方法认真计算海流流速的要求。可是,过去沿用至今的动力计算^[1],方法虽简便,但只能计算因密度分布所生的梯度流(或地转流),且存在着既不考虑风力,又不顾及湍流摩擦力,再加无运动面难以确定,即令设法作出浅海订正,其结果又往往与事实不符等根本性缺陷;而如籍 Ekman 漂流理论^[2]计算海流,又仅能计算因风所生的漂流,且还存在着既不考虑海水密度分布,又视海面无倾斜,再加湍流动力粘滞系数难以确定等与实际相差较远的理论依据。近代兴起的一些海流数值计算,又往往都局限于全流或深度平均流速的计算。因此,建立一种既考虑到海洋内部海水分布,又考虑到海面风力外加海面大气压力作用,顾及到海洋中湍流摩擦力,又体现流速随深度变化,而更重要的是应用起来简易的计算海流流速的理论及方法,便成为很需要解决的问题了。为此,本文作了这样的初步尝试,曾应用在东中国的海黑潮流系的流速计算中而获得了初步的成功。

在海流基本上是由海洋内部密度分布及海面风力和气压作用下形成的认识下,我们便可在水平压强梯度力、地转偏向力及由铅直湍流所生的水平湍流摩擦力的平衡下,依据海流动力学的运动方程、连续方程以及海面升降必须遵从的条件,求得由海水密度、风力及气压等决定的流速解析解,籍此将海洋内部各点的海流流速计算出来。而在求解的过程中,我们还提出了依据一定事实,确定水平压强梯度力及铅直湍流动力粘滞系数的理论及方法。当然本文所提出的计算海流流速的理论及方法,还存在着一些需待今后继续研究克服的缺点,例如仅考虑铅直湍流,适用于较深的海洋,需要预先知道海面流速等。

一、定常恒速海流流速计算

(一) 定常恒速海流流速计算

在忽略海水水平湍流摩擦力,且视海水流动为定常恒速时,考虑在水平压强梯度力、地转偏向力及由铅直湍流所在的水平湍流摩擦力三者取得平衡下的水平及铅直向运动方程将为:

$$-\frac{1}{\rho} \frac{\partial P}{\partial x} + 2\Omega \sin\varphi \cdot v + \frac{A_z}{\rho} \frac{\partial^2 u}{\partial z^2} = 0, \quad (1)$$

$$-\frac{1}{\rho} \frac{\partial P}{\partial y} - 2\Omega \sin\varphi \cdot u + \frac{A_z}{\rho} \frac{\partial^2 v}{\partial z^2} = 0, \quad (2)$$

$$-\frac{1}{\rho} \frac{\partial p}{\partial z} + g = 0. \quad (3)$$

连续方程为:

$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0. \quad (4)$$

式中 x 指向东, y 指向北, z 指向下; u 、 v 及 w 分别为流速的东分量、北分量及铅直向下分量; ρ 为海水密度; P 为海水压强; φ 为地理纬度; Ω 为地转角速度; g 为重力加速度; A_z 为铅直湍流动力粘滞系数。

引进水平复速度 W 为:

$$W = u + iv. \quad (5)$$

将 (1) 及 (2) 式合并为:

$$\frac{\partial^2 W}{\partial z^2} - [(1+i)a]^2 W = \frac{1}{A_z} \left(\frac{\partial P}{\partial x} + i \frac{\partial P}{\partial y} \right), \quad (6)$$

式中 a 为与 z 无关的参量, 即:

$$a = \sqrt{\frac{\rho \Omega \sin\varphi}{A_z}}. \quad (7)$$

求解 (6) 及 (4) 式的边界条件分别为:

$$\text{在海面 } z = -\zeta \text{ 处} \quad \left(-A_z \frac{\partial W}{\partial z} \right)_{-\zeta} = T_x + iT_y, \quad (8)$$

在海底或某一深度 $z = H$ 处

$$(W)_H = 0, \quad (9)$$

$$(\omega)_H = 0. \quad (10)$$

式中 $-\zeta$ 为海面的升高。 T_x 及 T_y 分别为风应力的东及北分量。

运用拉普拉斯变换求解 (6) 式。考虑到边界条件 (8), 令:

$$LW(x, y, z) = \tilde{W}(x, y, \lambda) = \tilde{W},$$

$$L \frac{\partial^2 W(x, y, z)}{\partial z^2} = \lambda^2 \tilde{W}(x, y, \lambda) - \lambda W_{-\zeta}(x, y, -\zeta) + \frac{1}{A_z} (T_x + iT_y)$$

$$= \lambda^2 \tilde{W} - \lambda W_{-\zeta} + \frac{1}{A_z} (T_x + iT_y),$$

$$L \left[\frac{\partial P(x, y, z)}{\partial x} + i \frac{\partial P(x, y, z)}{\partial y} \right] = \frac{\partial \tilde{P}(x, y, \lambda)}{\partial x} + i \frac{\partial \tilde{P}(x, y, \lambda)}{\partial y} = \frac{\partial \tilde{P}}{\partial x} + i \frac{\partial \tilde{P}}{\partial y},$$

则

$$\begin{aligned} \tilde{W} = & \frac{\lambda}{\lambda^2 - [(1+i)a]^2} W_{-\zeta} - \frac{1}{\lambda^2 - [(1+i)a]^2} \frac{T_x + iT_y}{A_z} \\ & + \frac{1}{\lambda^2 - [(1+i)a]^2} \frac{1}{A_z} \left(\frac{\partial \tilde{P}}{\partial X} + i \frac{\partial \tilde{P}}{\partial Y} \right). \end{aligned}$$

反变换后便求得包含海面水平流速 $W_{-\zeta}$ 的水平流速解 W 为:

$$\begin{aligned} W = & W_{-\zeta} \text{ch}(1+i)az - \frac{T_x + iT_y}{(1+i)aA_z} \text{sh}(1+i)az \\ & + \frac{1}{(1+i)aA_z} \int_{-\zeta}^z \left(\frac{\partial P}{\partial x} + i \frac{\partial P}{\partial y} \right)_z, \text{sh}(1+i)a(z-z') dz'. \quad (11) \end{aligned}$$

引进边界条件 (9) 便可求得海面流速解 $W_{-\zeta}$ 为:

$$\begin{aligned} W_{-\zeta} = & \frac{T_x + iT_y}{(1+i)aA_z} \frac{\text{sh}(1+i)aH}{\text{ch}(1+i)aH} \\ & - \frac{1}{(1+i)aA_z \text{ch}(1+i)aH} \int_{-\zeta}^H \left(\frac{\partial P}{\partial x} + i \frac{\partial P}{\partial y} \right)_z, \text{sh}(1+i)a(H-z') dz', \quad (12) \end{aligned}$$

解中第一部分为风力生成的, 第二部分则为由密度分布生成的流动, 将此式代入 (11) 式, 便可求得在任意深度的水平流速解为:

$$\begin{aligned} W = & \frac{T_x + iT_y}{(1+i)aA_z} \frac{\text{sh}(1+i)a(H-z)}{\text{ch}(1+i)aH} \\ & + \frac{1}{(1+i)aA_z} \int_{-\zeta}^z \left(\frac{\partial P}{\partial x} + i \frac{\partial P}{\partial y} \right)_z, \text{sh}(1+i)a(z-z') dz' \\ & - \frac{\text{ch}(1+i)az}{(1+i)aA_z \text{ch}(1+i)aH} \int_{-\zeta}^H \left(\frac{\partial P}{\partial x} + i \frac{\partial P}{\partial y} \right)_z, \text{sh}(1+i)a(H-z') dz', \quad (13) \end{aligned}$$

解中第一部分由风力生成的, 第二部分及第三部分则为由于密度分布生成的流动.

考虑到第三个运动方程, 即 (3) 式, 分部积分解 (12) 及 (13) 的积分项, 便得另一种形式的海面水平流速 $W_{-\zeta}$ 为:

$$\begin{aligned} W_{-\zeta} = & \frac{T_x + iT_y}{(1+i)aA_z} \frac{\text{sh}(1+i)aH}{\text{ch}(1+i)aH} - \frac{1}{[(1+i)a]^2 A_z} \left(\frac{\partial P}{\partial x} + \frac{\partial P}{\partial y} \right)_{-\zeta} \\ & + \frac{1}{[(1+i)a]^2 A_z \text{ch}[(1+i)aH]} \left(\frac{\partial P}{\partial x} + i \frac{\partial P}{\partial y} \right)_H \\ & - \frac{g}{[(1+i)a]^2 A_z \text{ch}(1+i)aH} \int_{-\zeta}^H \left(\frac{\partial P}{\partial x} + i \frac{\partial P}{\partial y} \right)_z, \text{ch}(1+i)a(H-z') dz', \quad (14) \end{aligned}$$

及另一种形式的在任意深处的水平流速 W 为:

$$W = \frac{T_x + iT_y}{(1+i)aA_z} \frac{\text{sh}(1+i)a(H-z)}{\text{ch}(1+i)aH} - \frac{1}{[(1+i)a]^2 A_z} \left(\frac{\partial P}{\partial x} + i \frac{\partial P}{\partial y} \right)_z,$$

$$\begin{aligned}
& + \frac{1}{[(1+i)a]^2 A_z} \frac{\operatorname{ch}(1+i)az}{\operatorname{ch}(1+i)aH} \left(\frac{\partial P}{\partial x} + i \frac{\partial P}{\partial y} \right)_H \\
& + \frac{g}{[(1+i)a]^2 A_z} \int_{-\zeta}^z \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} \operatorname{ch}(1+i)a(z-z') dz' \\
& - \frac{g}{[(1+i)a]^2 A_z} \frac{\operatorname{ch}(1+i)az}{\operatorname{ch}(1+i)aH} \int_{-\zeta}^H \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} \operatorname{ch}(1+i)a(H-z') dz'. \quad (15)
\end{aligned}$$

一旦水平压强梯度 $\left(\frac{\partial P}{\partial x} + i \frac{\partial P}{\partial y} \right)$ 在海面 $z = -\zeta$ 处, 在海底或某一深 $z = H$ 处以及在任意深度 z 处的值, 以及铅直湍流动力粘滞系数 A_z 的值能确定, 我们即可根据 (14) 及 (15) 式将在海面的和在任意深度的水平流速计算出。

(二) 水平压强梯度的确定

按照第三个运动方程, 即 (3) 式, 可知在任意深度 z 处

$$\begin{aligned}
\left(\frac{\partial P}{\partial x} + i \frac{\partial P}{\partial y} \right)_z & = g \int_{-\zeta}^z \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} dz' \\
& + \left[g \rho_{-\zeta} \left(\frac{\partial \xi}{\partial x} + i \frac{\partial \xi}{\partial y} \right) + \left(\frac{\partial P_a}{\partial x} + i \frac{\partial P_a}{\partial y} \right) \right]. \quad (16)
\end{aligned}$$

式中 P_a 为作用在海面的大气压强。

如将 (16) 式代入以水平复速度 W 表示的水平运动方程 (6), 得:

$$\frac{\partial^2 W}{\partial z^2} - [(1+i)a]^2 W = \frac{g}{A_z} \int_{-\zeta}^z \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} dz' + K, \quad (17)$$

式中

$$K = \frac{1}{A_z} \left[g \rho_{-\zeta} \left(\frac{\partial \xi}{\partial x} + i \frac{\partial \xi}{\partial y} \right) + \left(\frac{\partial P_a}{\partial x} + i \frac{\partial P_a}{\partial y} \right) \right]. \quad (18)$$

求解 (17) 式的边界条件仍为 (8) 及 (9) 式。

令水平流速 W 作如下代换, 即:

$$W = W' - \frac{K}{[(1+i)a]^2}, \quad (19)$$

代入 (17) 式, 则可改写为:

$$\frac{\partial^2 W'}{\partial z^2} - [(1+i)a]^2 W' = \frac{g}{A_z} \int_{-\zeta}^z \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} dz', \quad (20)$$

而边界条件 (8) 及 (9) 式将改为:

$$\text{在海面 } z = -\zeta \text{ 处, } \left(\frac{\partial W'}{\partial z} \right)_{-\zeta} = -\frac{T_x + iT_y}{A_z}, \quad (21)$$

$$\text{在海底或某一深 } z = H \text{ 处, } (W')_H = \frac{K}{[(1+i)a]^2}. \quad (22)$$

再运用拉普斯变换求解 (20) 式. 考虑到边界条件 (21) 式, 令:

$$\begin{aligned}
 LW'(x, y, z) &= \tilde{W}'(x, y, \lambda) = \tilde{W}', \\
 L \frac{\partial^2 W'(x, y, z)}{\partial z^2} &= \lambda^2 \tilde{W}'(x, y, \lambda) - \lambda W'_{-z}(x, y, -\xi) \\
 &+ \frac{1}{A_z} (T_x + iT_y) = \lambda^2 \tilde{W}' - \lambda W'_{-z} + \frac{1}{A_z} (T_x + iT_y) \\
 L \left\{ \int_{-z}^z \left[\frac{\partial \rho(x, y, z)}{\partial x} + i \frac{\partial \rho(x, y, z)}{\partial y} \right] dz' \right\} \\
 &= \int_{-z}^z \left[\frac{\tilde{\partial \rho}(x, y, \lambda)}{\partial x} + i \frac{\tilde{\partial \rho}(x, y, \lambda)}{\partial y} \right] d\lambda = \int_{-z}^z \left(\frac{\tilde{\partial \rho}}{\partial x} + i \frac{\tilde{\partial \rho}}{\partial y} \right)_{\lambda} d\lambda,
 \end{aligned}$$

反变换后便可求得包含海面水平流速 W'_{-z} 的解为 W' 为:

$$\begin{aligned}
 W' &= W'_{-z} \operatorname{ch}(1+i)az - \frac{T_x + iT_y}{(1+i)aA_z} \operatorname{sh}(1+i)az \\
 &+ \frac{g}{(1+i)aA_z} \int_{-z}^z \int_{-z}^{z'} \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_z \operatorname{sh}(1+i)a(z-z') dz'' dz'. \quad (23)
 \end{aligned}$$

引进边界条件(22), 便可求得海面水平流速解 W'_{-z} 为:

$$\begin{aligned}
 W'_{-z} &= \frac{T_x + iT_y}{(1+i)aA_z} \frac{\operatorname{sh}(1+i)aH}{\operatorname{ch}(1+i)aH} \\
 &- \frac{g}{(1+i)aA_z \operatorname{ch}(1+i)aH} \int_{-z}^H \int_{-z}^{z'} \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_z \operatorname{sh}(1+i)a \\
 &\times (H-z') dz'' dz' + \frac{K}{[(1+i)a]^2} \frac{1}{\operatorname{ch}(1+i)aH}. \quad (24)
 \end{aligned}$$

将此解代入 (23) 式, 便可求得水平流速解 W' 为:

$$\begin{aligned}
 W' &= \frac{T_x + iT_y}{(1+i)aA_z} \frac{\operatorname{sh}(1+i)a(H-z)}{\operatorname{ch}(1+i)aH} \\
 &+ \frac{g}{(1+i)aA_z} + \int_{-z}^z \int_{-z}^{z'} \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_z \operatorname{sh}(1+i)a(z-z') dz'' dz' \\
 &- \frac{g}{(1+i)aA_z} \frac{\operatorname{ch}(1+i)az}{\operatorname{ch}(1+i)aH} \int_{-z}^H \int_{-z}^{z'} \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_z \operatorname{sh}(1+i)a \\
 &\times (H-z') dz'' dz' + \frac{K}{[(1+i)a]^2} \frac{\operatorname{ch}(1+i)az}{\operatorname{ch}(1+i)aH}. \quad (25)
 \end{aligned}$$

再将之代入 (19) 式, 便可最后求得水平流速解 W 为:

$$W = \frac{T_x + iT_y}{(1+i)aA_z} \frac{\operatorname{sh}(1+i)a(H-z)}{\operatorname{ch}(1+i)aH}$$

$$\begin{aligned}
& + \frac{g}{(1+i)aA_z} \int_{-\zeta}^z \int_{-\zeta}^z \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z''} \text{sh}(1+i)a(z-z') dz'' dz' \\
& - \frac{g}{(1+i)aA_z} \frac{\text{ch}(1+i)az}{\text{ch}(1+i)aH} \int_{-\zeta}^H \int_{-\zeta}^{z'} \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} \text{sh}(1+i) \\
& \times a(H-z') dz'' dz' - \frac{K}{[(1+i)a]^2} \left[1 - \frac{\text{ch}(1+i)az}{\text{ch}(1+i)aH} \right]. \quad (26)
\end{aligned}$$

因而可求得海面水平流速解 $W_{-\zeta}$ 为:

$$\begin{aligned}
W_{-\zeta} &= \frac{T_x + iT_y}{(1+i)aA_z} \frac{\text{sh}(1+i)aH}{\text{ch}(1+i)aH} \\
& - \frac{g}{(1+i)aA_z} \frac{1}{\text{ch}(1+i)aH} \int_{-\zeta}^H \int_{-\zeta}^{z'} \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} \\
& \times \text{sh}(1+i)a(H-z') dz'' dz' - \frac{K}{[(1+i)a]^2} \frac{\text{ch}(1+i)aH - 1}{\text{ch}(1+i)aH}. \quad (27)
\end{aligned}$$

从 (27) 式我们可求得:

$$\begin{aligned}
- \frac{K}{[(1+i)a]^2} &= \frac{\text{ch}(1+i)aH}{\text{ch}(1+i)aH - 1} W_{-\zeta} - \frac{T_x + iT_y}{(1+i)aA_z} \frac{\text{sh}(1+i)aH}{\text{ch}(1+i)a - 1} \\
& + \frac{g}{(1+i)aA_z} \frac{1}{\text{ch}(1+i)aH - 1} \int_{-\zeta}^H \int_{-\zeta}^{z'} \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} \\
& \times \text{sh}(1+i)a(H-z') dz'' dz'. \quad (28)
\end{aligned}$$

将这样求得的 $-\frac{K}{[(1+i)a]^2}$ 代入 (26) 式, 便得:

$$\begin{aligned}
W &= \frac{\text{ch}(1+i)aH - \text{ch}(1+i)az}{\text{ch}(1+i)aH - 1} W_{-\zeta} \\
& + \frac{\text{sh}(1+i)a(H-z)[\text{ch}(1+i)aH - 1] - \text{sh}(1+i)aH[\text{ch}(1+i)aH - \text{ch}(1+i)az]}{(1+i)aA_z \text{ch}(1+i)aH[\text{ch}(1+i)aH - 1]} \\
& \times (T_x + iT_y) + \frac{g}{(1+i)aA_z} \int_{-\zeta}^z \int_{-\zeta}^{z'} \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} \text{sh}(1+i)a(z-z') dz'' dz' \\
& + \frac{[1 - \text{ch}(1+i)az]}{(1+i)aA_z [\text{ch}(1+i)aH - 1]} \int_{-\zeta}^H \int_{-\zeta}^{z'} \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} \text{sh}(1+i) \\
& \times a(H-z') dz'' dz'. \quad (29)
\end{aligned}$$

解中第一项为海面流速贡献部分, 第二项为风生海流部分, 第三项及第四项均为密度分布所生的海流部分。

如对 (29) 式求 z 的二阶导数, 得:

$$\frac{\partial^2 W}{\partial z^2} = - \frac{[(1+i)a]^2 \text{ch}(1+i)az}{\text{ch}(1+i)aH - 1} W_{-\zeta}$$

$$\begin{aligned}
& + \frac{(1+i)a\{\text{sh}(1+i)a(H-z)[\text{ch}(1+i)aH-1] + \text{sh}(1+i)aH \cdot \text{ch}(1+i)az\}}{A_z \text{ch}(1+i)aH[\text{ch}(1+i)aH-1]} \\
& \times (T_x + iT_y) + g \frac{(1+i)a}{A_z} \int_{-\zeta}^z \int_{-\zeta}^{z'} \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} \text{sh}(1+i)a(z-z') dz'' dz' \\
& + \frac{g}{A_z} \int_{-\zeta}^z \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} dz' - g \frac{(1+i)a}{A_z} \frac{\text{ch}(1+i)az}{\text{ch}(1+i)aH-1} \\
& \times \int_{-\zeta}^H \int_{-\zeta}^{z'} \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} \text{sh}(1+i)a(H-z') dz'' dz'. \quad (30)
\end{aligned}$$

经过分部积分后得:

$$\begin{aligned}
\frac{\partial^2 W}{\partial z^2} & = - \frac{[(1+i)a]^2 \text{ch}(1+i)az}{\text{ch}(1+i)aH-1} W_{-\zeta} \\
& + \frac{(1+i)a\{\text{sh}(1+i)a(H-z)[\text{ch}(1+i)aH-1] + \text{sh}(1+i)aH \text{ch}(1+i)az\}}{A_z \text{ch}(1+i)aH[\text{ch}(1+i)aH-1]} \\
& \times (T_x + iT_y) + \frac{g}{A_z} \int_{-\zeta}^z \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} \text{ch}(1+i)a(z-z') dz' \\
& + \frac{g}{A_z} \frac{\text{ch}(1+i)az}{\text{ch}(1+i)aH-1} \left[\int_{-\zeta}^H \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} dz' \right. \\
& \left. - \int_{-\zeta}^H \left(\frac{\partial \rho}{\partial x} + \frac{\partial \rho}{\partial y} \right)_{z'} \text{ch}(1+i)a(H-z') dz' \right]. \quad (31)
\end{aligned}$$

而在海底或某一深 $z = H$ 处

$$\begin{aligned}
\left(\frac{\partial^2 W}{\partial z^2} \right)_H & = - \frac{[(1+i)a]^2 \text{ch}(1+i)aH}{\text{ch}(1+i)aH-1} W_{-\zeta} + \frac{[(1+i)a] \text{sh}(1+i)aH}{A_z [\text{ch}(1+i)aH-1]} (T_x + iT_y) \\
& + \frac{g}{A_z} \frac{\text{ch}(1+i)aH}{\text{ch}(1+i)aH-1} \int_{-\zeta}^H \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} dz' \\
& - \frac{g}{A_z} \frac{1}{\text{ch}(1+i)aH-1} \int_{-\zeta}^H \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} \text{ch}(1+i)a(H-z') dz'. \quad (32)
\end{aligned}$$

由 (6) 式, 知在 $z = H$ 处

$$\left(\frac{\partial^2 W}{\partial z^2} \right)_H = \frac{1}{A_z} \left(\frac{\partial P}{\partial x} + i \frac{\partial P}{\partial y} \right)_H, \quad (33)$$

将之代入 (32) 式, 便得:

$$\begin{aligned}
\frac{1}{A_z} \left(\frac{\partial P}{\partial x} + i \frac{\partial P}{\partial y} \right)_H & = - \frac{[(1+i)a]^2 \text{ch}(1+i)aH}{\text{ch}(1+i)aH-1} W_{-\zeta} \\
& + \frac{[(1+i)a] \text{sh}(1+i)aH}{A_z [\text{ch}(1+i)aH-1]} (T_x + iT_y)
\end{aligned}$$

$$\begin{aligned}
& + \frac{g}{A_z} \frac{\operatorname{ch}(1+i)aH}{\operatorname{ch}(1+i)aH-1} \int_{-\zeta}^H \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} dz' \\
& - \frac{g}{A_z} \frac{1}{\operatorname{ch}(1+i)aH-1} \int_{-\zeta}^H \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} \operatorname{ch}(1+i)a(H-z') dz', \quad (34)
\end{aligned}$$

稍经整理便得:

$$\begin{aligned}
& \frac{1}{[(1+i)a]^2 A_z} \frac{\operatorname{ch}(1+i)aH-1}{\operatorname{ch}(1+i)aH} \left(\frac{\partial P}{\partial x} + i \frac{\partial P}{\partial y} \right)_H \\
& = -W_{-\zeta} + \frac{1}{(1+i)aA_z} \frac{\operatorname{sh}(1+i)aH}{\operatorname{ch}(1+i)aH} (T_x + iT_y) \\
& + \frac{g}{[(1+i)a]^2 A_z} \int_{-\zeta}^H \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} dz' \\
& - \frac{1}{[(1+i)a]^2 A_z} \frac{1}{\operatorname{ch}(1+i)aH} \int_{-\zeta}^H \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} \operatorname{ch}(1+i)a(H-z') dz'. \quad (35)
\end{aligned}$$

在 $aH \geq 3$ 的情况下, 我们即可将上式右边近似地改写为:

$$\begin{aligned}
& \frac{1}{[(1+i)a]^2 A_z} \left(\frac{\partial P}{\partial x} + i \frac{\partial P}{\partial y} \right)_H = -W_{-\zeta} + \frac{1}{(1+i)aA_z} \frac{\operatorname{sh}(1+i)aH}{\operatorname{ch}(1+i)aH} (T_x + iT_y) \\
& + \frac{g}{[(1+i)a]^2 A_z} \int_{-\zeta}^H \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} dz' \\
& - \frac{g}{[(1+i)a]^2 A_z} \frac{1}{\operatorname{ch}(1+i)aH} \int_{-\zeta}^H \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} \operatorname{ch}(1+i)a(H-z') dz'. \quad (36)
\end{aligned}$$

由于按照 (16) 式, 知在海面 $z = -\zeta$ 处

$$\left(\frac{\partial P}{\partial x} + i \frac{\partial P}{\partial y} \right)_{-\zeta} = g\rho_{-\zeta} \left(\frac{\partial \zeta}{\partial x} + i \frac{\partial \zeta}{\partial y} \right) + \left(\frac{\partial P_a}{\partial x} + i \frac{\partial P_a}{\partial y} \right), \quad (37)$$

以及在海底或某一深度 $z = H$ 处,

$$\left(\frac{\partial P}{\partial x} + i \frac{\partial P}{\partial y} \right)_H = g \int_{-\zeta}^H \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} dz' + \left(\frac{\partial P}{\partial x} + i \frac{\partial P}{\partial y} \right)_{-\zeta}. \quad (38)$$

于是再将 (38) 式代入 (36) 式, 便可近似地求得:

$$\begin{aligned}
& \frac{1}{[(1+i)a]^2 A_z} \left(\frac{\partial P}{\partial x} + i \frac{\partial P}{\partial y} \right)_{-\zeta} = -W_{-\zeta} + \frac{1}{(1+i)aA_z} \frac{\operatorname{sh}(1+i)aH}{\operatorname{ch}(1+i)aH} (T_x + iT_y) \\
& - \frac{g}{[(1+i)a]^2 A_z} \frac{1}{\operatorname{ch}(1+i)aH} \int_{-\zeta}^H \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_{z'} \operatorname{ch}(1+i)a(H-z') dz'. \quad (39)
\end{aligned}$$

而如将因此获得的 (39) 式与 (14) 式相比较, 即可近似地求得:

$$\left(\frac{\partial P}{\partial x} + i \frac{\partial P}{\partial y} \right)_H = 0. \quad (40)$$

将 (40) 式代入 (37) 式, 可近似地求得在海面 $z = -\xi$ 处,

$$\left(\frac{\partial P}{\partial x} + i \frac{\partial P}{\partial y} \right)_{-\xi} = -g \int_{-\xi}^H \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_z dz', \quad (41)$$

而如再将 (41) 式及 (37) 式代入 (16) 式, 即可近似地求得在任意深处

$$\left(\frac{\partial P}{\partial x} + i \frac{\partial P}{\partial y} \right)_z = -g \int_z^H \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_z dz'. \quad (42)$$

于是, 解析解中有关水平压强梯度项便可用水平密度梯度项来代替.

(三) 水平流速及 A_z 的确定

在水平压强梯度如上述那样确定以后, 我们便可将它们代入 (14) 及 (15) 式中, 去求得海面水平流速 $W_{-\xi}$ 及在任意深处的水平流速 W 分别为:

$$\begin{aligned} W_{-\xi} = & \frac{T_x + iT_y}{(1+i)aA_z} \frac{\text{sh}(+i)aH}{\text{ch}(+i)aH} + \frac{g}{[(1+i)a]^2 A_z} \int_{-\xi}^H \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_z dz' \\ & - \frac{g}{[(1+i)a]^2 A_z \text{ch}(1+i)aH} \int_{-\xi}^H \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_z \text{ch}(1+i)a(H-z') dz', \quad (43) \end{aligned}$$

及

$$\begin{aligned} W = & \frac{T_x + iT_y}{(1+i)aA_z} \frac{\text{sh}(1+i)a(H-z)}{\text{ch}(1+i)aH} + \frac{g}{[(1+i)a]^2 A_z} \int_z^H \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_z dz' \\ & + \frac{g}{[(1+i)a]^2 A_z} \int_{-\xi}^z \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_z \text{ch}(1+i)a(z-z') dz' \\ & \times \frac{g}{[(1+i)a]^2 A_z} \frac{\text{ch}(1+i)az}{\text{ch}(1+i)aH} \int_{-\xi}^H \left(\frac{\partial \rho}{\partial x} + i \frac{\partial \rho}{\partial y} \right)_z \text{ch}(1+i)a(H-z') dz'. \quad (44) \end{aligned}$$

如将 (43) 及 (44) 写成分量式, 便为:

$$\begin{aligned} u_{-\xi} = & \frac{(T_x + T_y)\text{sh}2aH - (T_y - T_x)\text{sin}2aH}{2aA_z(\text{ch}2aH + \cos2aH)} + \frac{g}{2a^2 A_z} \int_{-\xi}^H \left(\frac{\partial \rho}{\partial y} \right)_z dz' \\ & + \frac{g}{2a^2 A_z(\text{ch}2aH + \cos2aH)} \int_{-\xi}^H \left(\frac{\partial \rho}{\partial x} \right)_z [\text{sh}a(2H-z')\text{sin}az' \\ & + \text{sin}az' \text{sin}a(2H-z')] dz' - \frac{g}{2a^2 A_z(\text{ch}2aH + \cos2aH)} \int_{-\xi}^H \left(\frac{\partial \rho}{\partial y} \right)_z \\ & \times [\text{cha}(2H-z')\text{cos}az' + \text{ch}az'\text{cosa}(2H-z')] dz', \quad (45) \\ v_{-\xi} = & \frac{(T_y - T_x)\text{sh}2aH + (T_x + T_y)\text{sin}2aH}{2aA_z(\text{ch}2aH + \cos2aH)} - \frac{g}{2a^2 A_z} \int_{-\xi}^H \left(\frac{\partial \rho}{\partial x} \right)_z dz' \end{aligned}$$

$$\begin{aligned}
& + \frac{g}{2a^2 A_z (\text{ch}2aH + \cos2aH)} \int_{-\zeta}^H \left(\frac{\partial \rho}{\partial x} \right)_z, [\text{cha}(2H - z') \cos az' \\
& + \text{chaz}' \cos a(2H - z')] dz' + \frac{g}{2a^2 A_z (\text{ch}2aH + \cos2aH)} \int_{-\zeta}^H \left(\frac{\partial \rho}{\partial y} \right)_z, \\
& \times [\text{sha}(2H - z') \sin az' + \text{shaz}' \sin a(2H - z')] dz', \quad (46)
\end{aligned}$$

及

$$\begin{aligned}
u_z = & \frac{(T_x + T_y) [\text{sha}(2H - z) \cos az - \text{shaz} \cos a(2H - z)]}{2a A_z (\text{ch}2aH + \cos2aH)} \\
& + \frac{(T_y - T_x) [\text{cha}(2H - z) \sin az - \text{chaz} \sin a(2H - z)]}{2a A_z (\text{ch}2aH + \cos2aH)} \\
& + \frac{g}{2a^2 A_z} \int_z^H \left(\frac{\partial \rho}{\partial y} \right)_z, dz' + \frac{g}{2a^2 A_z} \int_{-\zeta}^z \left[\left(\frac{\partial \rho}{\partial y} \right)_z, \text{cha}(z - z') \cos a(z - z') \right. \\
& + \left. \left(\frac{\partial \rho}{\partial x} \right)_z, \text{cha}(z - z') \sin a(z - z') \right] dz' \\
& + \frac{g}{a^2 A_z} \frac{\text{cha}H \cos aH \text{chaz} \cos az + \text{sha}H \sin aH \text{shaz} \sin az}{\text{ch}2aH + \cos2aH} \\
& \times \int_{-\zeta}^H \left[\left(\frac{\partial \rho}{\partial y} \right)_z, \text{cha}(H - z') \cos a(H - z') + \left(\frac{\partial \rho}{\partial x} \right)_z, \text{sha}(H - z') \right. \\
& \times \left. \sin a(H - z') \right] dz' + \frac{g}{a^2 A_z} \frac{\text{sha}H \sin aH \text{chaz} \cos az - \text{cha}H \cos aH \text{shaz} \sin az}{\text{ch}2aH + \cos2aH} \\
& \times \int_{-\zeta}^H \left[\left(\frac{\partial \rho}{\partial x} \right)_z, \text{cha}(H - z') \cos a(H - z') - \left(\frac{\partial \rho}{\partial y} \right)_z, \text{sha}(H - z') \right. \\
& \times \left. \sin a(H - z') \right] dz', \quad (47)
\end{aligned}$$

$$\begin{aligned}
v_z = & \frac{(T_y - T_x) [\text{sha}(2H - z) \cos az - \text{shaz} \cos a(2H - z)]}{2a A_z (\text{ch}2aH + \cos2aH)} \\
& - \frac{(T_x + T_y) [\text{cha}(2H - z) \sin az - \text{chaz} \sin a(2H - z)]}{2a A_z (\text{ch}2aH + \cos2aH)} \\
& - \frac{g}{2a^2 A_z} \int_z^H \left(\frac{\partial \rho}{\partial x} \right)_z, dz' - \frac{g}{2a^2 A_z} \int_{-\zeta}^z \left[\left(\frac{\partial \rho}{\partial x} \right)_z, \right. \\
& \times \left. \text{cha}(z - z') \cos a(z - z') - \left(\frac{\partial \rho}{\partial y} \right)_z, \text{sha}(z - z') \sin a(z - z') \right] dz' \\
& + \frac{g}{a^2 A_z} \frac{\text{cha}H \cos aH \text{chaz} \cos az + \text{sha}H \sin aH \text{shaz} \sin az}{\text{ch}2aH + \cos2aH} \\
& \times \int_{-\zeta}^H \left[\left(\frac{\partial \rho}{\partial x} \right)_z, \text{cha}(H - z') \cos a(H - z') - \left(\frac{\partial \rho}{\partial y} \right)_z, \text{sha}(H - z') \sin a(H - z') \right] dz'
\end{aligned}$$

$$\begin{aligned}
 & -z')dz' + \frac{g}{a^2 A_z} \frac{\text{sha}H\text{sin}aH\text{ch}az\text{cos}az - \text{cha}H\text{cos}aH\text{sh}az\text{sin}az}{\text{ch}2aH + \text{cos}2aH} \\
 & \times \int_{-z}^H \left[\left(\frac{\partial \rho}{\partial y} \right)_{z'} \text{cha}(H-z')\text{cos}a(H-z') + \left(\frac{\partial \rho}{\partial x} \right)_{z'} \text{sha}(H-z') \right. \\
 & \left. \times \text{sin}a(H-z') \right] dz'. \tag{48}
 \end{aligned}$$

将一系列 A_z 代入 (45) 及 (46) 式, 使这样计算出的 u_{-z} 及 v_{-z} 与实测的海面流速相比较, 使差值小于某个预定的可允许数值, 即可确定所要计算海区海面上各点的 A_z 值. 再将这样确定的 A_z 值代入 (47) 及 (48) 式, 即可计算出在任意深度处的流速.

我们便是这样先确定东中国海黑潮流系各点的 A_z , 而后计算得出各点的水平流速的.

这样确定 A_z 的方法, 其缺点是要在预知海面水平流速的情况下, 始能比较作出, 这在海面水平流速未知的情况下, 就无法使用该法确定 A_z 了.

至于海流流速的铅直分量, 在水平流速分量已求得的情况下, 则可将连续方程 (4) 先化作差分方程而后进行计算, 计算时可认为海底的铅直流速 $\omega_H = 0$.

(四) 东中国海黑潮流系的海流流速计算

我们利用二月份风应力和密度值, 对界于北纬 $24^\circ 15' - 34^\circ$, 东经 $120^\circ 30' - 130^\circ 15'$ 东中国海的黑潮流系的海流流速进行计算, 计算是在每点的 A_z 均定为 375 克/厘米·秒的情况进行的. 计算时采用东西步长为半个经度, 南北步长为半个纬度和铅直步长按水深标准层次间隔设置的立体网格. 网格上放置密度数据, 网格中心放置风应力数据及所要计算的流速值. 上层次四个水平流速计算值和下层次四个水平流速值之间的中心点, 放置所要计算的铅直流速值. 利用差分法进行计算, H 最大值为 800 米. 计算结果表明:

1. 黑潮进入东中国海后其主要分支为: (1) 流入台湾海峡的向南逆流约在北纬 $25^\circ 15' - 26^\circ 15'$, 东经 $121^\circ 30' - 123^\circ 15'$ 附近流入台湾海峡. (2) 闽浙支流约在北纬 $28^\circ 30' - 29^\circ 45'$, 东经 $124^\circ 45' - 126^\circ 15'$ 成几乎平行于闽浙海岸的方向向北流去. (3) 约在北纬 $29^\circ 55' - 30^\circ 15'$, 东经 $128^\circ 45' - 130^\circ$ 处流向日本南部成为黑潮延续体. (4) 黄海暖流约在北纬 $30^\circ 45' - 32^\circ 45'$, 东经 $126^\circ 15' - 127^\circ 15'$ 附近朝西北方向流入黄海.

(5) 对马暖流大约在北纬 $29^\circ 57' - 33^\circ 45'$, 东经 $120^\circ 15' - 129^\circ 30'$ 附近北上.

2. 黑潮的主轴是波状, 这一主轴的位置, 几乎与中国海 200 米等深线平行, 开始偏于 200 米等深线东南方向的 50 海里, 而后北上, 直到北纬 30° 处便折向东. 由此可见, 黑潮主干自我国台湾北上后, 大致是沿大陆架和大陆坡毗连区域流动. 如果我们在这个波状主轴之间作一中线, 则可得黑潮主轴波动的振幅约为 62 海里, 波长约为 440 海里. 从南面往下到 500 米, 主轴的位置由于受大陆架的挤压, 虽然略偏东, 但偏离的并不大.

3. 黑潮流系中出现三处较大流速值, 分别位于北纬 $30^\circ 15'$ 、东经 129° , 北纬 $26^\circ 45'$ 、东经 $126^\circ - 126^\circ 15'$ 和北纬 $24^\circ 45'$ 、东经 123° 三处, 该三处的深度各为 800 米, 2000 米和 1200 米. 出现较大流速处的主要原因是由于地形的影响. 或是由于地形逐渐变窄, 黑潮大部分从这里流进太平洋; 或是由于当黑潮北上往东偏时, 受到冲绳岛的阻挡; 或是由于正当台湾与西表岛间, 黑潮水大量涌进东海所致.

4. 在黑潮的右侧, 因岛屿和地形的摩擦作用, 常发生很强的逆流, 约在北纬 $27^{\circ}15'$ 、东经 $127^{\circ}45'$ 和北纬 $25^{\circ}30'$ — $27^{\circ}30'$, 东经 $121^{\circ}30'$ — 124° 处。

我们对中国海海区的黑潮流系, 进行了初步的海流流速计算。利用实测与计算海面流速的比较来确定铅直湍流动力粘滞系数 A_z 是合理的, 据此进一步计算各点的流速是可行的, 从而获得东中国海二月份黑潮流轴的空间分布及流系流速的空间分布。

由于计算所利用的资料是不同年代的, 非准同时的, 而且实测海面流速尚待进一步完善, 致使 A_z 无法按细分的小海区来选取, 以示其随经、纬度的变化。计算结果与实测海面流速相比有些偏小。其它季节的资料正准备进一步计算, 以便研究黑潮的季节变化。

在我们的计算过程中, 多承北京京字 116 部队的冯康同志及金旦华同志在编制计算程序中诚挚的指导和帮助, 在此致以衷心感谢。又多承中国科学院海洋研究所及国家海洋局情报研究所提供有关资料, 特在此一并致以诚挚的谢意。

参 考 文 献

- (1) Sandström, J.M., B.Helland-Hansen, Über Berechnung von Meereströmungen, *Reports on Norwegian Fishery and Marine Investigations*, 2(1903), 4.
- (2) Ekman, V.W., On the influence of the earth's rotation on ocean currents, *Arkiv Mat. Astron. Fysik*, 2(1905), 11.

STUDY ON THE CALCULATION OF OCEAN CURRENT VELOCITY

Jing Zhenhua

(Shandong College of Oceanology)

ABSTRACT

In order to calculate the wind induced current and the density current which vary with ocean depth the author presents a method for calculating the current velocity from the distributions of wind stresses at the sea surface and the density distribution of sea water. The horizontal gradients of pressure at any determined by making certain assumptions, and the coefficients of vertical eddy viscosity are obtained by comparing the calculated values of the current velocities at sea surface with the observed ones. We used this method to calculate the current velocities

at any depth in Kuroshio, since it gives comparatively satisfactory results in the positions of the branches of Kuroshio, the situations of the current axis of Kuroshio etc., the validity of this method is verified.